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PU	Public	X
PP	Restricted to other programme participants (including the Commission Services)	
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Abstract

Internal loading has a fundamental role in the nutrient recycling and ecosystem functioning of large and shallow lakes, but it can also influence deep lakes. This deliverable includes one article and a manuscript, that deal with the temperature-induced changes on internal loading in shallow and deep lakes, and long-term trends and seasonal changes in P retention in shallow Danish lakes.

Manuscript by Järvinen et al.

Marko Järvinen, Kari Kallio & Ahti Lepistö

Effects of temperature changes on internal loading in shallow and deep lakes

Manuscript

Impacts of temperature changes on the internal loading of nutrients in shallow and deep lakes were assessed using high-frequency monitoring data from a shallow eutrophic lake basin in southern Finland and output from literature. Temperature affects internal loading largely by influencing stability of the water column, and the duration of the ice-covered period. Resultant responses to the duration and intensity of stratification and mixing influence the primary drivers of internal nutrient loading in the water column and the water-sediment interface: dissolved oxygen concentrations, Redox-conditions, temperature-dependent metabolism and microbial degradation, and resuspension. Higher temperature is likely to lead to longer and stronger stratification episodes in summer which may increase the importance of O₂ and Redox conditions in the phosphorus (P) release relative to resuspension in relatively shallow lakes. In deep lakes, longer stratification results in accumulation of nutrients in the hypolimnion and the nutrients typically become available to upper water layers only after degradation of thermal stratification in autumn or by partial breakdown of the thermocline. This together with expected delayed freezing dates in future can strengthen and prolong the duration of autumnal diatom and cyanobacterial blooms also in deep lakes. Daily data from Säkylän Pyhäjärvi support the finding that shallow lakes react rapidly to changes in thermal regime, and mixing brings nutrients and other suspended material to surface water layers which will increase algal biomass and also the probability for cyanobacterial blooms.

Article by Søndergaard et al.

Martin Søndergaard, Rikke Bjerring & Erik Jeppesen

Persistent internal phosphorus loading during summer in shallow eutrophic lakes.

Hydrobiologia DOI 10.1007/s10750-012-1091-3 (Springer Online FirstArticles 2012)

Twentyone years of monthly mass balance and lake chemistry data from six shallow eutrophic and fast flushed eutrophic Danish lakes were used to investigate long-term trends in yearly and seasonal patterns of P retention. In only one of the lakes, the external P input has experienced major changes during the last >20 years. The lakes showed a distinct seasonal pattern with high P concentrations and typically negative P retention during summer (up to -300% of the external loading from May to August). During winter, P retention was up to 50% of the external loading from December to April. Internal P loading from the sediment delayed lake recovery by approximately 10 years in the lake with the most recently reduced external loading, but in all the lakes net release of P from the sediment occurred during summer. P release in the lakes has not abated during the past decade, indicating that the sediment of eutrophic and turbid shallow lakes remains a net source of P during summer. The seasonal variations in P retention became more pronounced with increasing P levels, and retention decreased with increasing temperature, but increased if clear water conditions were established.

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Manuscript

EFFECTS OF TEMPERATURE CHANGES ON INTERNAL LOADING IN SHALLOW AND DEEP LAKES

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29 **Abstract**

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31 Impacts of temperature changes on the internal loading of nutrients in shallow and deep lakes were
32 assessed using high-frequency monitoring data from a shallow eutrophic lake basin in southern
33 Finland and output from literature. Temperature affects internal loading largely by influencing
34 stability of the water column, and the duration of the ice-covered period. Resultant responses to the
35 duration and intensity of stratification and mixing influence the primary drivers of internal nutrient
36 loading in the water column and the water-sediment interface: dissolved oxygen concentrations,
37 Redox-conditions, temperature-dependent metabolism and microbial degradation, and resuspension.
38 Higher temperature is likely to lead to longer and stronger stratification episodes in summer which
39 may increase the importance of O₂ and Redox conditions in the phosphorus (P) release relative to
40 resuspension in relatively shallow lakes. In deep lakes, longer stratification results in accumulation
41 of nutrients in the hypolimnion and the nutrients typically become available to upper water layers
42 only after degradation of thermal stratification in autumn or by partial breakdown of the
43 thermocline. This together with expected delayed freezing dates in future can strengthen and
44 prolong the duration of autumnal diatom and cyanobacterial blooms also in deep lakes. Daily data
45 from Säkylän Pyhäjärvi support the finding that shallow lakes react rapidly to changes in thermal
46 regime, and mixing brings nutrients and other suspended material to surface water layers which will
47 increase algal biomass and also the probability for cyanobacterial blooms.

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51 **Keywords**

52 resuspension, mixing, thermal stratification, phosphorus, climate change

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54 **Introduction**

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56 Internal loading of nutrients has a fundamental role in the nutrient recycling of large and shallow
57 lakes and it is often the reason for the failure or delay of successful lake management of eutrophic
58 (shallow) lakes, by supporting blooms of diatoms and harmful cyanobacteria, also after reductions
59 in external nutrient loading (e.g. Søndergaard et al. 2003, 2012, Håkanson 2004, French &
60 Petticrew 2007). The ultimate reason behind internal loading is generally a long eutrophication
61 history: external loading has led to high lake nutrient levels, increased phytoplankton biomass and
62 increased sedimentation of organic matter to the bottom areas of the lake, which eventually leads to
63 phosphorus (P) release from the P-rich surface sediments (e.g. Kleeberg & Kozerski 1997, Ventelä
64 et al. 2011).

65

66 Internal loading of nutrients in lakes depends on several mechanisms and in-lake processes that are
67 often coupled to each other: water temperature, resuspension, Redox conditions, pH, iron (Fe)
68 chemistry, chemical diffusion, bioturbation, mineralization and microbial processes in the
69 sediments and the presence of macrophytic vegetation in the littoral and shallow lake areas (e.g.
70 Boström et al. 1982, Søndergaard et al. 2003, Niemistö 2008). Many of the listed mechanisms have
71 a strong dependency on the meteorological forcing and they are thus sensitive to any changes in
72 future climate (French & Petticrew 2007, Adrian et al. 2009, Pettersson et al. 2010).

73

74 Air temperature, wind speed, cloud cover and turbulent kinetic energy are the principal drivers of
75 thermal stability of the water column that is expressed as the depth of the thermocline, and changes
76 in thermal stratification and mixing of lakes (Imboden et al. 1983, Adrian et al. 2009). Shallow
77 lakes typically have weak thermal stability which increases the likelihood of intermittent
78 breakdowns of stratification during the open water periods even under moderate wind conditions.

79

80 Effects of climate change on lakes are complex (e.g. Blenckner 2001, George 2010). Responses are
81 modified by regional differences, and they strongly depend on the characteristics of the catchment
82 and lake mixing regimes (Blenckner 2001, Nöges et al. 2009, Adrian et al. 2009). It is evident that
83 internal loading has a key influence on the nutrient availability and ecosystem productivity in
84 eutrophic shallow lakes, and frequently mixing polymictic lakes (Kleeberg & Kozerski 1997,
85 Søndergaard et al. 2003, 2012). Due to shallowness and small water volume, shallow lakes are also
86 likely more susceptible to direct impacts of changes in temperature. Internal loading caused by
87 mixing events and internal wave activity can also be important in the shallow and isolated parts or

88 in the littoral of deep dimictic lakes, and obviously in the pelagial during spring and autumn
89 overturns (e.g. Niemistö 2008).

90

91 Here we discuss on how temperature-induced changes affect internal nutrient loading in lakes
92 with varying depth by using results from a shallow Finnish polymictic lake and output from
93 literature. We demonstrate that the internal loading has importance throughout the summer in
94 shallow lakes, whereas in the pelagial of dimictic deep lakes the impacts of internal loading are
95 generally discernible following the late summer-autumn break-down of stratification. As the
96 temperature increase accelerates many mechanisms of internal loading (e.g. Nöges et al. 2011a), we
97 suggest that in the future warmer climate, P release in shallow lakes is likely to increase. Although
98 macrophyte communities play an important role in the recycling of nutrients and overall the
99 ecosystem functioning of shallow lakes (Jeppesen et al. 1997, Søndergaard et al. 2003) their
100 responses to temperature changes and implications to internal loading will not be dealt here.

101

102 **Methods**

103

104 The study is largely based on high-frequency (one hour measurement interval) monitoring data
105 from May-October 2009 from the shallow and polymictic lake Pyhäjärvi (Säkylä, SW Finland.
106 Pyhäjärvi is a large (155 km²) and shallow ($z_m = 5.5$ m) lake suffering from eutrophication.
107 Increased eutrophication of Pyhäjärvi has been a major concern since the late 1980s as blooms of
108 cyanobacteria became more frequent particularly during the 1990s. Between 1970 and 1992 total P
109 (TP) concentrations doubled, but have decreased since then (Ventelä et al., 2007). Recently, new
110 information of spatial variability of cyanobacterial biomass in different parts of the lake has been
111 obtained from moving-boat transect measurements (Lepistö et al., 2010). The main concerns for the
112 status of the lake, also with respect to changing climate, are possible changes in external nutrient
113 loads (Ekholm et al., 1997, Ventelä et al., 2011).

114

115 For Pyhäjärvi, hourly vertical profiles of water temperature and oxygen were recorded at the depths
116 of 5, 10 and 19 m by a real-time lake raft which located in the deepest part of lake in May-October
117 2009 (Kallio et al., 2010). Turbidity, chlorophyll *a* fluorescence (phytoplankton biomass) and
118 phycocyanin fluorescence (cyanobacteria biomass) have been recorded hourly during the same
119 period at the depth of 1 m. Water chemistry data (see below for details) have been collected
120 manually from the depths of 1, 5, 10, 15, 20 and 24 m in 1980-2008 and in 2009-2011 from the

121 depths of 1, 15 and 24 m. The results from the depths of 0-5 and 15-24 m were averaged to
122 represent values in the epi- and hypolimnion, respectively.

123

124 Phosphate phosphorus (PO₄-P), total phosphorus (TP), ammonium nitrogen (NH₄-N) and iron (Fe)
125 were analyzed using standard methods (e.g. APHA, 2000). Stability of the water column was
126 calculated as the Schmidt stability (J m⁻²) from the daily water temperature profiles. The
127 stratification periods were distinguished from the mixing/weak stratification periods using the
128 stability criterium of >3.5 J m⁻², which corresponded well with the water temperature difference
129 exceeding 1 °C between the surface and the bottom layer.

130

131 For the synthesis we also compiled results from previously published studies dealing with internal
132 loading and its mechanisms as well as its relationships with the meteorological forcing and the
133 future climate change.

134

135 **Results and discussion**

136

137 Temperature affects internal loading indirectly by influencing the stability conditions of the water
138 column (Soranno et al. 1997, Wilhelm & Adrian 2008, Niemistö et al. 2007, 2009). High-resolution
139 measurements of water temperature show that stability conditions of Säkylän Pyhäjärvi changed
140 rapidly during the open water season of 2009 (Fig. 1). Stratification collapsed several times during
141 the measurement period, being strongest from late June till early July. The periods of low stability
142 were discernible as increased deep water concentrations of O₂, and turbidity of the uppermost water
143 layers, both indicating increased mixing of the water column (Fig. 1). Despite intermittent mixing
144 periods during summer, which temporarily improved hypolimnetic O₂ concentrations, deep water
145 O₂ decreased throughout the summer from mid-June until the start of autumn turn-over in mid-
146 August. As a whole, O₂ and P concentrations in the hypolimnion differed between the contrasting
147 stability conditions in Pyhäjärvi in 2009-2010. O₂ concentrations were significantly lower (p=0.006,
148 Mann-Whitney U-test, n=12) and PO₄-P concentrations higher (p=0.004, n=12) in the hypolimnion
149 during periods of stratification (Fig. 2; see also Fig.1). The same was observed for ammonium-
150 nitrogen (NH₄-N; p=0.002) where hypolimnetic concentrations were on average 4.5-times higher
151 (180 µg l⁻¹, n=5) during stratified periods than during mixed periods (40 µg l⁻¹, n=7) in June-
152 September. Total-P concentrations were at the same level (p=0.328 epi; p=0.082 hypo, n=12)
153 among the different stability periods, and also the epilimnetic PO₄-P did not differ significantly
154 between the periods (p=0.195, n=12) (Fig. 2). There are also large annual differences in the summer

155 O₂ and nutrient concentrations in Pyhäjärvi (Fig. 3). Deep water (23-25 m) average O₂
156 concentrations are typically lowest and inorganic N and P concentrations highest in late winter
157 under the ice in 1980-2011. The annual variation in PO₄-P and NH₄-N in the deep water layer is
158 also highest during winter. Ice-out in April-May is followed by a complete overturn and the whole
159 water column becomes oxygenated and this with the nutrient uptake by spring phytoplankton could
160 explain the decrease of PO₄-P and NH₄-N concentrations also in the hypolimnion (Fig. 3). During
161 summer (June-September) O₂ and nutrient concentrations show high variation between the years
162 depending largely on the prevailing stratification and mixing events. The results of Pyhäjärvi do not
163 allow detection of mechanisms behind the above described patterns in internal nutrient loading in
164 summer, but they are likely related to changes in Redox conditions as well as resuspension (see e.g.
165 Søndergaard et al. 2003). The release of P from sediments often results from Redox-dependent
166 release of P bound to iron and microbial processes which can also happen in shallow well-mixed
167 eutrophic lakes with organic rich sediments because the oxic layer of the sediment cannot prevent
168 release of P from the deeper parts of the sediment (Søndergaard et al. 2003). Fe concentrations were
169 highest near bottom under the ice during winter, and the Fe concentrations and their variation
170 between the years were clearly lower in summer, when Pyhäjärvi was mixing several times.
171 Although resuspended nutrients are generally mostly in particulate form (Ekholm et al. 1997),
172 resuspension has been shown to generally increase both total phosphorus (P) and inorganic P
173 concentrations in the water (Niemistö et al. 2008, French & Petticrew 2007, Søndergaard et al.
174 2011). The change in total P could not be found in Pyhäjärvi (Fig. 2), possibly due the rather limited
175 number of observations.

176
177 As also indicated by Pyhäjärvi data, temperature can influence directly and indirectly on many key
178 variables responsible for internal loading (Fig. 4, Søndergaard et al. 2003). Mineralization of
179 organic material and the metabolism rates of microbes in the sediment depend on water temperature
180 (e.g. Boström et al. 1982). As a result of frequent mixing of the water column during summer which
181 is typical for many shallow lakes, warmer surface water is transported and mixed to deeper water
182 layers, thus heating the hypolimnion and the upper layers of the sediment. Higher water
183 temperatures in the hypolimnion will accelerate processes in the settled organic material, O₂
184 consumption and release of P (Boström et al. 1982, Koski-Vähälä et al. 2011, Niemistö 2008).
185 Søndergaard et al. (2012) showed that the seasonal P retention in 21 eutrophic shallow Danish lakes
186 was negatively related to water temperature in summer when higher temperature and a thin oxidized
187 surface sediment layer probably resulted in a Redox-sensitive P release from the sediments. The
188 changes in hypolimnetic temperature are less likely in deep stratified lakes where both increases and

189 decreases have been reported in relation to meteorological forcing (Arvola et al. 2010, Jones et al.
190 2010).

191

192 Temperature affects directly also the duration of ice-covered period in the lakes of northern Europe
193 (Blenckner et al. 2004). Higher air water temperatures cause earlier ice-break in spring and later
194 freezing in autumn, thus increasing the length of the open water period (Fig. 4, Elo et al. 1998,
195 Blenckner et al. 2010). In deep lakes this may lead, together with higher surface water temperatures,
196 to longer stratification (Blenckner et al. 2010, Jones et al. 2010), which would increase the duration
197 of isolation of deeper water layers and would thus increase the probability for O₂ depletion,
198 decrease in Redox conditions and the dissolution of P from the sediments in eutrophic, but also in
199 humic lakes during summer (Fig. 4, Forsius et al. 2010). However, this does not necessarily lead to
200 significant changes in P concentrations during autumn mixing as with the advent of autumnal
201 circulation, ferrous iron (Fe²⁺) is rapidly oxidized and a large proportion of the phosphate is
202 precipitated as ferric phosphate (e.g. Wetzel 2001). Despite this, increased mixing may have an
203 important role in the loading of P and for cyanobacterial blooms in summer by deepening of the
204 thermocline and consequent upwelling of nutrients (Soranno et al. 1997).

205

206 In shallow lakes, the prolongation of ice-free period can intensify resuspension and increase the
207 annual load of suspended solids and nutrients from the sediment. For instance, Niemistö et al.
208 (2007) estimated, using scenario model by Kuusisto (1989), that doubling of CO₂ would increase 28
209 % the annual loading of total P by resuspension in a shallow and eutrophic basin of Lake Hiidenvesi
210 in southern Finland, due to a shortening of ice-covered period and more intensive mixing during
211 open water period. For deeper Finnish lakes, PROBE model simulations have suggested
212 prolongation of summer stratification (Elo et al. 1998). Longer stratification periods in shallow
213 lakes would increase the importance of O₂ and Redox conditions relative to resuspension (Behrend
214 et al. 1993). Niemistö (2008) estimated that for some shallow lakes with greater than average
215 turbulence, large littoral areas and small anaerobic hypolimnia, P release from sediments can be
216 much larger than values (10-30% of total loading) for deeper, more stratified lakes.

217

218 As climate change is likely to lead higher water temperatures in summer (Fig. 4, e.g. Arvola et al.
219 2010) this will evidently favour the occurrence of large-sized cyanobacteria in eutrophic lakes
220 (Paerl & Huisman 2008, Wagner & Adrian 2009, Blenckner et al. 2010, Kosten et al. 2012), that are
221 less edible to grazing zooplankton. This will in turn increase the sedimentation of organic matter
222 and consumption of O₂ in the bottom layers of lakes. High phytoplankton production may also

223 increase pH to a level (pH >9) where P desorption may occur, in particular when the bottom water
 224 temperature is high enough, >15°C (Koski-Vähälä et al. 2001, Niemistö 2008, French & Petticrew
 225 2007).

226
 227 The lakes in Europe and North America show increasing trends in dissolved organic carbon (water
 228 colour), and this so-called brownification has been related to climate-induced changes in
 229 precipitation and DOC loading from the catchment to lakes and/or the recovery of lakes and their
 230 catchments from acidification (e.g. Hongve et al. 2004, Monteith et al. 2007). Possible increases
 231 in water colour by humic substances in future will influence thermal regime of lakes by affecting
 232 the light absorption and stratification characteristics (e.g. Forsius et al. 2010), which are likely
 233 reflected also in internal loading.

234

235 **Conclusions**

236

237 To conclude, temperature changes influence internal loading in lakes mainly by modifying the
 238 stability conditions of the water column, and influencing the duration of ice-covered period in
 239 northern Europe. The exchange of P between sediments and the overlying water depends on the
 240 ability of the sediments to retain P, the conditions of the overlying water and the biota within the
 241 sediments that modify exchange equilibria and affect P transport back to the water (Boström et al.
 242 1982, Søndergaard et al. 2003). Overall, the results from Pyhäjärvi and literature suggest that the
 243 reduction of nutrient loads in particular in shallow lakes suffering from eutrophication may prove
 244 difficult in future due to impacts of warmer temperatures that support internal loading.

245

246

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248

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446

447 **FIGURE CAPTIONS**

448

449 **Figure 1.** Daily variation of Schmidt stability ($J m^{-2}$) (grey bars), minimum dissolved oxygen
 450 concentration ($O_2 mg l^{-1}$) at 19 m depth (black solid line), and maximum turbidity (FNU) (dashed
 451 line) at 1 m depth, and the concentration of phosphate-P ($PO_4-P \mu g l^{-1}$) in the hypolimnion (black
 452 dots) of Säkylän Pyhäjärvi (SW Finland) in May-October 2009.

453

454 **Figure 2.** Box plot presentation of Schmidt stability ($J m^{-2}$), dissolved O_2 concentration in
 455 hypolimnion ($mg l^{-1}$), phosphate-P ($PO_4-P \mu g l^{-1}$) and total phosphorus (tot-P $\mu g l^{-1}$) in the
 456 epilimnion (epi) and hypolimnion (hypo) of Lake Pyhäjärvi during stratified (str) and mixed (mix)
 457 periods in June-August 2009-2010. $n=5$ for str and $n=7$ for mix, except for stability where $n=37$ and
 458 $n=55$, respectively. Median values with 25th and 75th percentiles, and minimum and maximum
 459 values (whiskers). * = statistically significant difference ($p < 0.05$; Mann-Whitney U-test).

460

461 **Figure 3.** Mean monthly (1980-2011) concentrations of A) dissolved oxygen ($mg l^{-1}$), B) iron (Fe
 462 $\mu g l^{-1}$), C) phosphate-P ($PO_4-P \mu g l^{-1}$) and D) ammonium-nitrogen ($NH_4-N \mu g l^{-1}$) in the
 463 hypolimnion (23-25 m) of Lake Pyhäjärvi (SW Finland). Solid smoothed lines = average, broken
 464 smoothed lines = 5th and 95th percentiles. $n=5-67$ for different months.

465

466 **Figure 4.** A conceptual chart presenting the impacts of increased temperature on key mechanisms
 467 of internal loading in shallow lakes. The dashed line separates the epimnion and hypolimnion. P =
 468 phosphorus, Fe = iron, S = sulphur, OC = organic carbon, WT = water temperature, O_2 = dissolved
 469 oxygen.

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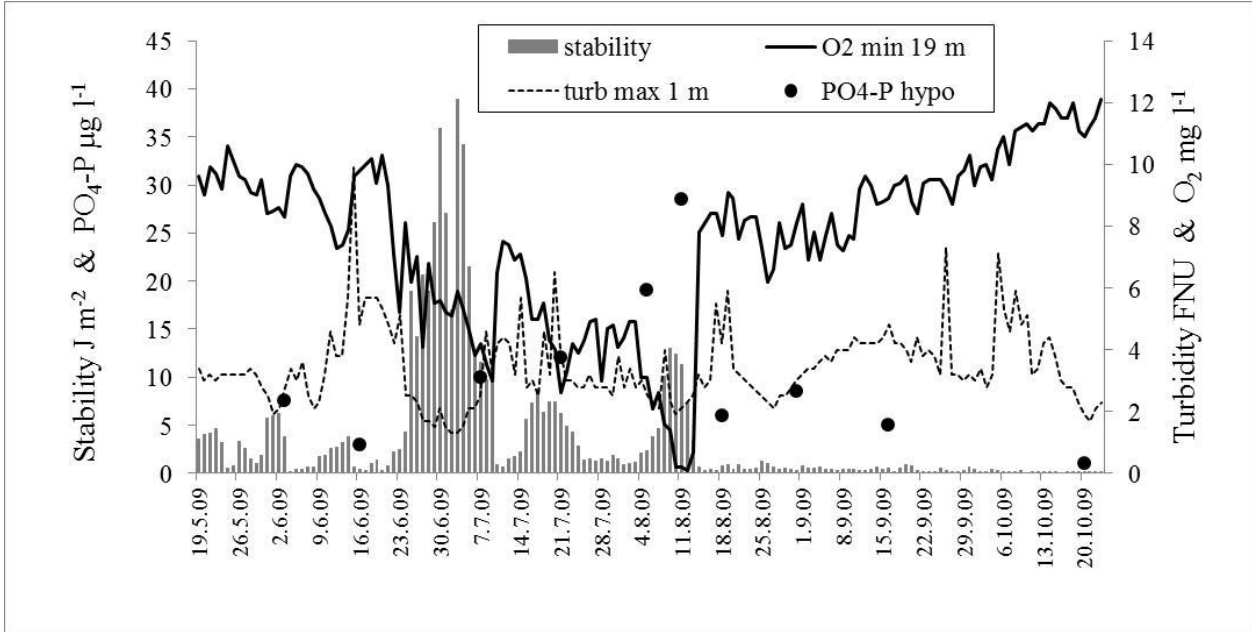
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472 **Fig 1.**

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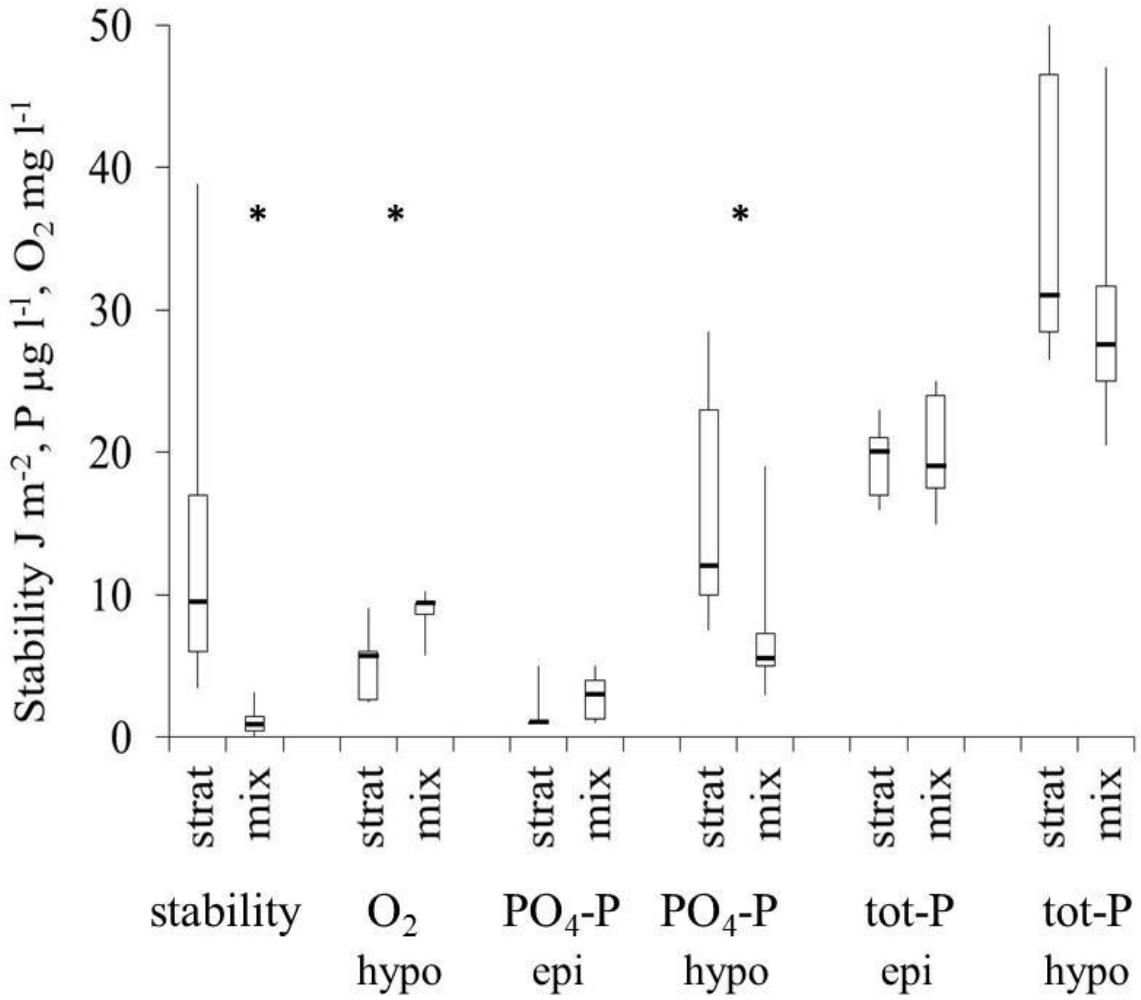
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478 **Fig. 2**

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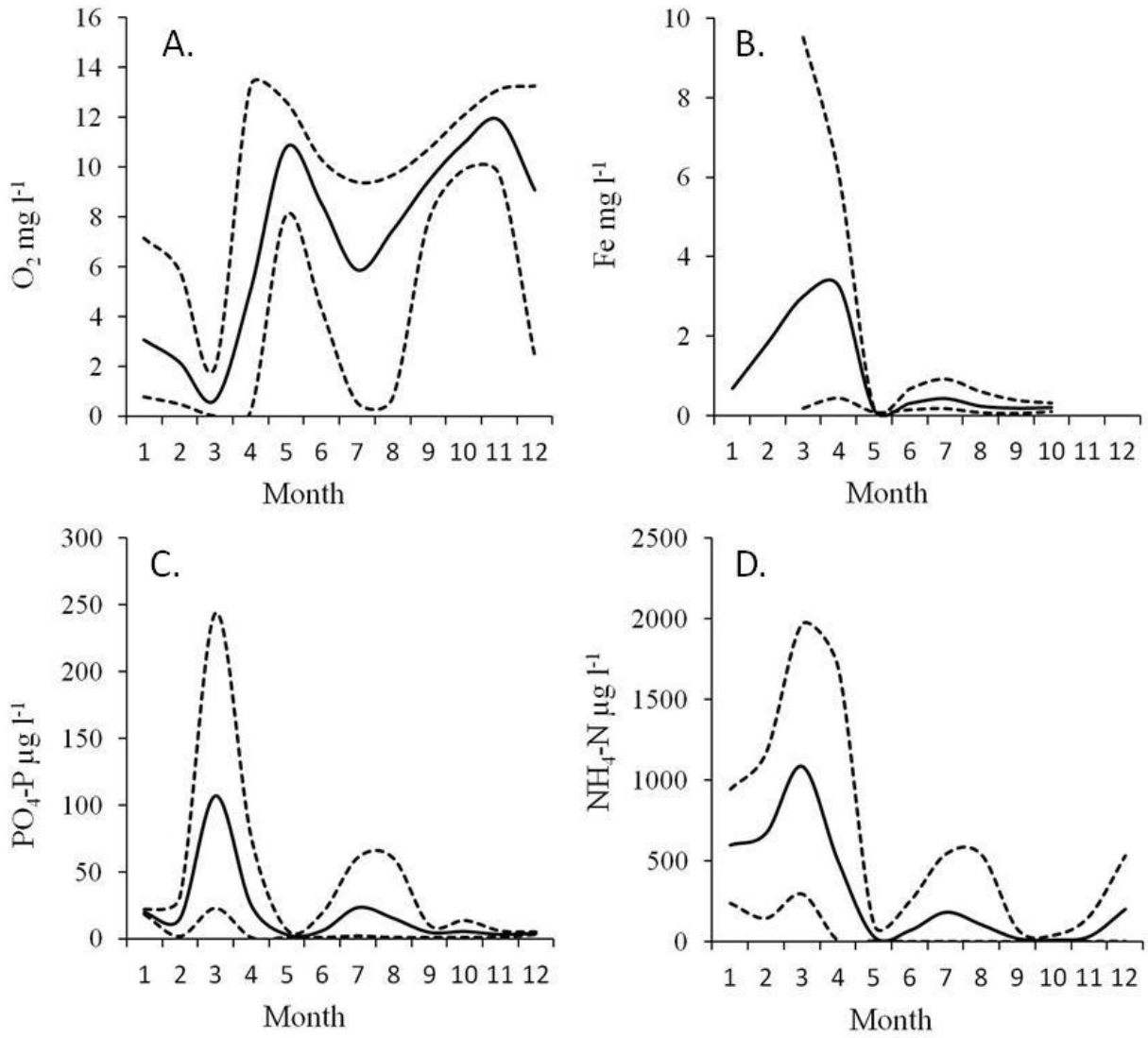
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485 **Fig. 3**

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Fig. 4

